

## SEISMIC EVIDENCE FOR BLIND THRUSTING OF THE NORTHWESTERN FLANK OF THE VENEZUELAN ANDES

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*Abstract.* Surface geology and seismic and well data from the northwestern flank of the Venezuelan Andes indicate overthrusting of Andean basement rocks toward the adjacent Maracaibo Basin along a blind thrust fault. The frontal monocline is interpreted as the forelimb of a northwestward verging fault-related fold deformed over a crustal-scale ramp. The Andean block has been thrust 20 km to the northwest and uplifted 10 km on a ramp that dips about 20°-30° southeastward. The thrust fault ramps up through crystalline basement rocks to a decollement horizon within the shaly units of the Cretaceous Colon-Mito Juan formations. Backthrusts in the monocline produce a wedge geometry and reduce the amount of blind slip required on the decollement northwest of the Andes. The rigid Andean uplift was caused by northwest-southeast compressive tectonic forces related to the convergence of the Caribbean plate, the Panama volcanic arc, and northwestern South America. The thick (up to 6 km) molasse deposits accumulated in the foredeep basin indicate that the Venezuelan Andes started to rise as early as the early Miocene. However, a late Miocene intramolasse unconformity marks the beginning of the formation of the monocline and the greatest uplift. The crustal-scale fault-related fold model may explain structural features seen in other areas of basement-involved foreland deformation.

### INTRODUCTION

Over the past several years, diverse tectonic models have been proposed to explain the Venezuelan Andes uplift. Bucher [1952], Gonzalez de Juana [1952], Hargraves and Shagam [1969], and Shagam [1972] described the Venezuelan Andes as a mega-anticline containing a horst-graben block complex and limited at both flanks by high-angle reverse faults. Rod [1960], Deratmiroff [1971], Schubert [1985], White [1985], Boesi et al. [1988], and Monsalve [1988] proposed that the Venezuelan Andes is a "mushroomlike" transpressional uplift with shear deformation in the Bocono fault system and imbricate thrusting toward both Andean fronts. Hospers and Van Wijnen [1959], Kellogg and Bonini [1982, 1985], Giegengack [1984], and Kohn et al. [1984] noted the asymmetry of the gravity field [Bellizia et al., 1976] and suggested that the Venezuelan Andes uplift was caused by crustal shortening along a southeast dipping "Wind River type" thrust fault.

This paper focuses on the structural interpretation of seismic sections traversing the northwestern flank of the Venezuelan Andes and the southern edge of the Maracaibo Basin between

La Fria and El Vigia (Figures 1 and 2). In this paper we present geological and geophysical evidence for basement-involved thrusting of the Venezuelan Andes toward the Maracaibo Basin.

### GEOLOGIC SETTING

The regional geologic setting of the area can be described as a mountain front adjacent to a foredeep basin. The mountain front represents a Precambrian-Paleozoic igneous-metamorphic complex partially covered with sediments of Mesozoic and Cenozoic ages. The foredeep is an asymmetric structural basin that contains up to 12 km of sedimentary section. The early Miocene to Pleistocene molasse deposits up to 6 km in thickness dominate the sedimentary section. The oldest sedimentary rocks that crop out in the mountain front are Jurassic red beds which are unconformably overlain by a Cretaceous sequence up to 2 km in thickness. From early Cretaceous to middle Eocene, sediments were deposited in an extensive passive continental margin setting (Figure 3). Coarse clastic sediments of the Rio Negro Formation, bioclastic carbonates of the Apon Formation, and siliciclastic sediments of the Aguardiente Formation were deposited first in the initial stage of the Cretaceous transgressive event. Limestones, calcareous shales, and black shales of the Capacho and La Luna formations represent anoxic facies deposited during the maximum marine transgression of late Cretaceous (Cenomanian-Campanian) age. Regression began in the latest Cretaceous (Maestrichtian) with marine deposition of the shaly Colon Formation and the sandier Mito-Juan Formation. This regression event continued through the Paleocene and middle Eocene with deposition of sandstones, siltstones, shales, and coals of the Orocué Group in a deltaic environment followed by fluvial sandstones of the Mirador Formation. Unconformably overlying this is the upper Eocene paralic facies of the Carbonera Formation which is in turn transitionally covered by lacustrine shales of the Oligocene Leon Formation. Miocene to Recent aged sediments represent continental deposits (Guayabo Group) derived mainly from erosion of the Venezuelan Andes uplift.

The sediments in the mountain front are folded into a steep northwestward dipping monocline. The dip of strata ranges from 25° to overturned, and the strike is northeast-southwest. Numerous faults, mainly oriented north northwest-south southeast and northeast-southwest, cut and offset the outcropping stratigraphy. The Rio Mocoties fault is considered to be the southwestern continuation of the Bocono right-lateral strike-slip fault.

The gravity field of the area shows a Bouguer anomaly in the north Andean trough of nearly -150 mGal [Bonini et al., 1977]. This large negative anomaly in the southern edge of the Maracaibo Basin indicates that the basin and the Andes are not in local isostatic equilibrium.

### Tectonic Setting

The intra-Miocene unconformities indicate a major Andean orogeny which has been related to the collision of the Panama island arc with the South American continent [Mann and Burke, 1984; Eva et al., 1989]. The resulting collision and the oblique subduction of the Carnegie ridge beneath the Andes in Ecuador may have caused the uplift of the Eastern Cordillera of Colombia and the detachment of the North Andes block or microplate along the Bocono and the East Andean frontal fault

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Paper number 93TC01893  
0278-7407/93/93TC-01893\$10.00

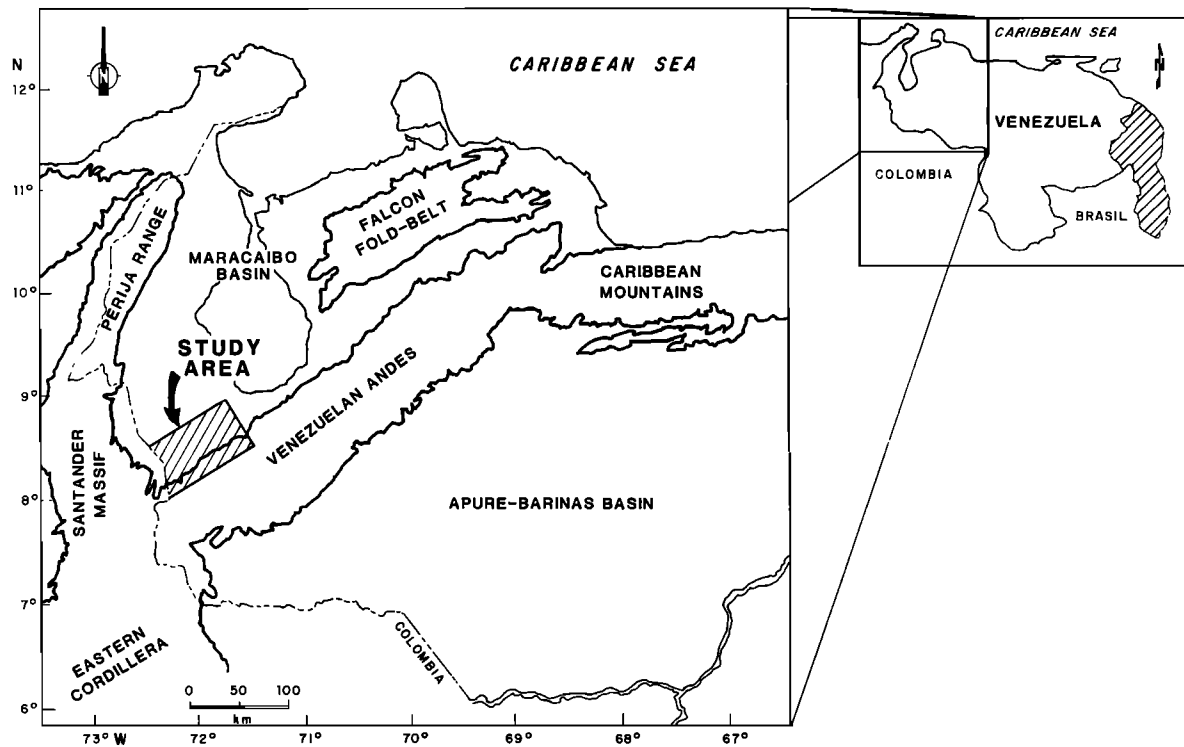


Fig. 1. Location map of the study area.

system [Pennington, 1981; Kellogg et al., 1985]. The limited cumulative displacement on the Bocono fault [Stephan, 1977; Gonzalez de Juana et al., 1980; Schubert, 1982; Giegengack, 1984], despite the relatively rapid present-day slip rate of 1 cm/yr [Schubert, 1980; Aggarwal, 1983], suggests that most of the northeastward movement of the North Andes block occurred in the last 5 or 10 m.y. [Kellogg, 1984; Macellari, 1984].

Global plate motion models predict west northwest-east southeast convergence for the Caribbean and South American plates [Jordan, 1975; Minster and Jordan, 1978; DeMets et al., 1990]. The north-south vector component is even greater for Caribbean-North Andean convergence when the North Andean movement northeastward relative to South America on the Bocono fault system is added. The direction of maximum shortening in the Venezuelan Andes is northwest-southeast. Kellogg and Bonini [1982, 1985] and Kellogg [1984] suggested that this vector of deformation is related to slow convergence ( $1.7 \pm 0.7$  cm/yr) between the Caribbean plate and the North Andean microplate at the South Caribbean marginal fault (Figure 4). This compressive deformation was also responsible for the uplifts of the Sierra de Perija and the Sierra Nevada de Santa Marta. Present-day amagmatic subduction along the South Caribbean marginal fault is indicated by active folding and thrusting in the South Caribbean deformed belt [Lehner et al., 1983; Ladd et al., 1984] and by a southeast dipping Wadati-Benioff zone [Dewey, 1972; Pennington, 1981; Kellogg and Bonini, 1982; Toto and Kellogg, 1992]. Pennington [1981] and Daniels [1991] estimated the dip of the Wadati-Benioff zone as  $20^\circ$ . Preliminary results from Central and South America (CASA) Global Positioning System (GPS) satellite geodetic measurements over a 3-year period

(1988-1991) show of the order of 1 to 2 cm/yr convergence between the Caribbean, Panama, and North Andean plates [Vega et al., 1991]. "Thick-skinned" basement-involved deformation has been correlated with magmatic gaps and low-angle or flat subduction in the Pampeanas Range of Argentina [Jordan et al., 1983; Jordan and Allmendinger, 1986] and the North American Cordillera [Lipman et al., 1971; Burchfiel and Davis, 1975; Coney, 1976; Dickinson and Snyder, 1978].

#### *Comparisons to Other Basement-Involved Foreland Deformations*

Structural features in the Venezuelan Andes are similar to those associated with the Laramide orogeny (40-70 m.y.) in the central and southern Rocky Mountains of the United States. The basement block tectonic style of the Laramide orogeny is characterized by the movement of basement blocks along reverse faults. The overlying sediments in many areas fold over reverse faulted basement blocks forming the monoclines [Stearns, 1971; Stearns and Weinberg, 1975; Matthews, 1978; Narr, 1993; Narr and Suppe, submitted manuscript] analogous to the bruchfalten (fault folds) of Stille [1925]. Local secondary normal faults occur [Wise, 1963; Howard, 1966; Stearns, 1971], but gravity data and Consortium for Continental Reflection Profiling (COCORP) deep seismic reflection profiling show that regional shortening far exceeds extension [Berg, 1962; Berg and Romberg, 1966; Smithson et al., 1978]. The structural relief on the basement surfaces exceeds 12 km for the Wind River and Sweetwater uplifts [Prucha et al., 1965].

Other active Andean mountain belts that are structurally analogous to the basement-involved Laramide orogeny in the

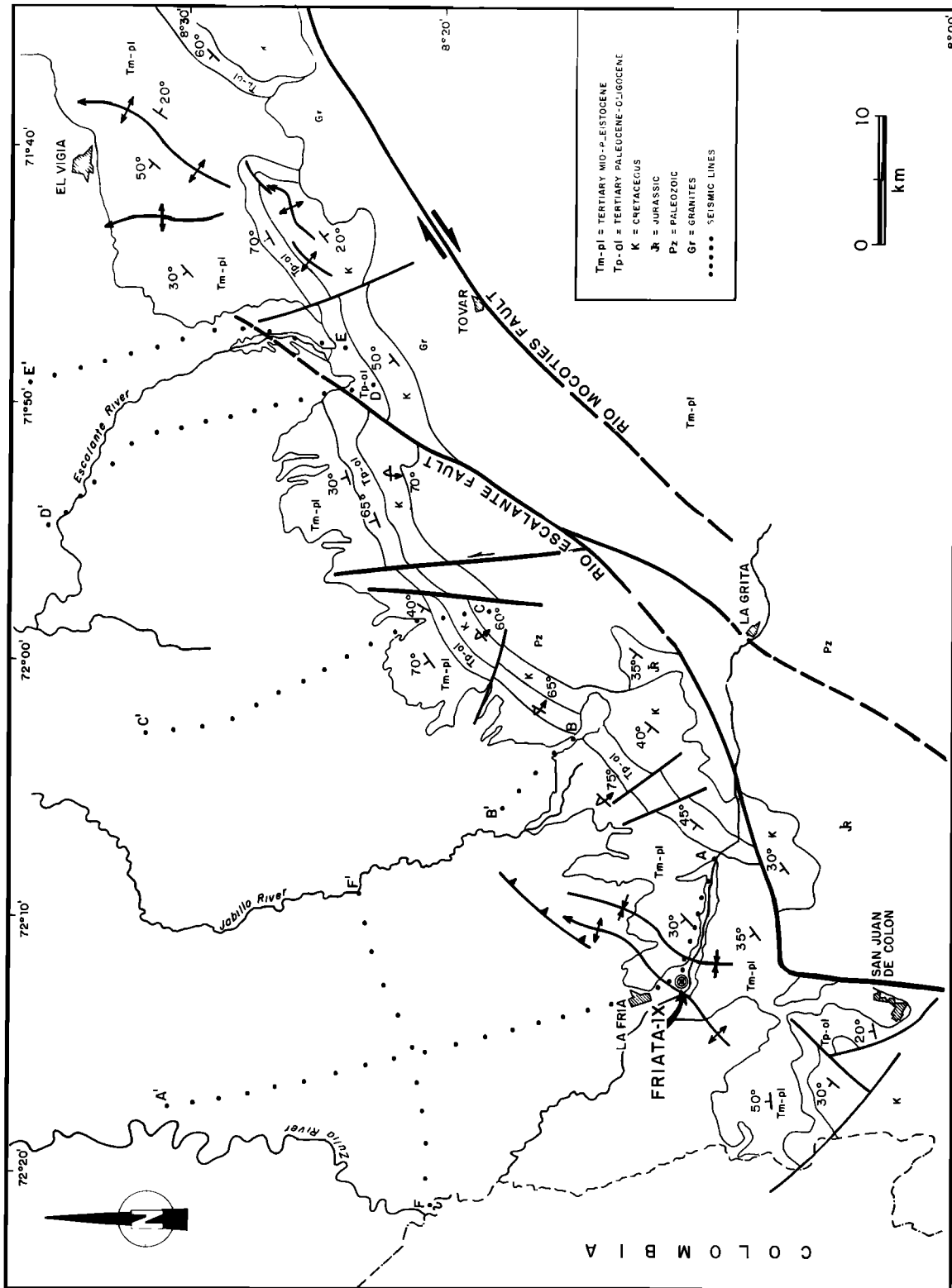


Fig. 2. Generalized geologic map of the study area and location of the seismic lines used in this study.

GEOLOGICAL TIME	AGE IN MA	STRATIGRAPHIC UNITS	DEPOSITIONAL ENVIROMENT	
PLEISTOCENE	2	BETIJOQUE	FLUVIAL ALLUVIAL	
PLIOCENE				
MIOCENE	5	EL GUAYABO GROUP	FLUVIO-DELTAIC	
	6			
	10			
OLIGOCENE	25	EL PALMAR	COASTAL PLAIN	
	36	LEON	PARALIC TO LAGOONAL	
EOCENE	U M L	CARBONERA		
		39	MIRADOR	FLUVIAL
		49		
PALEOCENE	54	OROCUE GROUP	PARALIC-DELTAIC	
CRETACEOUS	66	MITO JUAN	OPEN MARINE	
	74	COLON		
	MARINE	LA LUNA	OPEN MARINE ANOXIC	
		CAPACHO	CARBONATE PLATFORM DELTAIC	
		AGUARDIENTE		
	APON	FLUVIAL		
	RIO NEGRO			
PRE-CRETACEOUS	116	LA QUINTA		
		BASEMENT		

Fig. 3. Stratigraphy of the north Andean flank.

middle and southern Rocky Mountains include the Sierra de Perija on the Colombian-Venezuelan border and the Sierra Nevada de Santa Marta of Colombia [Kellogg and Bonini, 1982], the Peruvian Andes [Audebaud et al., 1973], and the Sierras Pampeanas of Argentina [Jordan et al., 1983; Jordan

and Allmendinger, 1986]. A similar style has been recognized in middle and northern Germany [Stille, 1910; Lotze, 1938], the Arbuckle-Wichita Mountains [Fichter, 1954], the Permian basin of west Texas [Elam, 1969], and the Laramide fault block uplifts and related monoclines of the Colorado Plateau [Kelly, 1955].

DATA AND METHOD

The data used in this study include seismic sections, surface geology, and the Friata-1X well log [DeToni, 1990]. Five seismic lines traversing the Andean mountain front between La Fria and El Vigia (Figure 2) were selected for this study.

To interpret the seismic profiles, the time sections were converted to depth sections: the travel times of prominent seismic horizons were converted to approximate depths in meters, using stacking velocities. Wherever possible, these seismic horizons were correlated with the exposed surface stratigraphy. The geologic solutions presented in this paper are qualitative descriptions and do not presume to describe the structure quantitatively (i.e., the cross sections are not balanced/retrodeformable).

SEISMIC INTERPRETATION

In seismic profiles of the area, two seismic provinces are recognized: the Andean frontal monocline and the foredeep basin. The southern portion which corresponds to the frontal monocline is characterized by discontinuous seismic horizons dipping to the northwest. The major seismic reflectors can be correlated with geologic units exposed at the surface. The northern portions of the seismic sections correspond to the foredeep basin. This zone shows well-defined, continuous seismic reflectors dipping gently to the southeast. The deeper reflectors are discontinuous at the mountain front and cannot be correlated with the reflectors present in the frontal

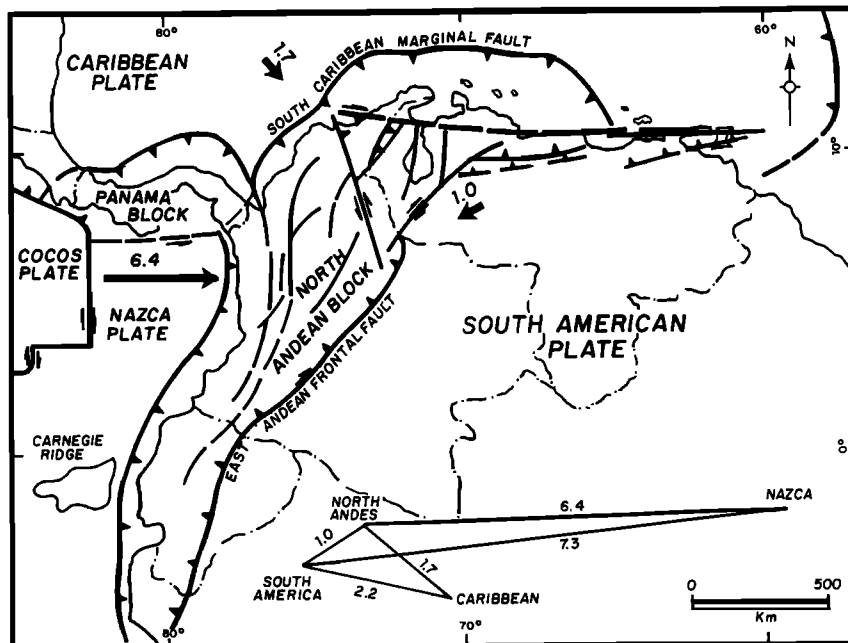


Fig. 4. Present-day plate motions (solid arrows) relative to the north Andean block showing average slip rates (in centimeters per year) during the last 5-10 m.y. [after Kellogg and Bonini, 1985].

monocline. As a result of this lack of lateral correlation, some Cretaceous units in the foredeep basin were identified by correlation of acoustic character with an interpreted seismic line and a synthetic trace for a well in the Los Manueles oil field (Figure 5).

#### Frontal Monocline

Mesozoic and Cenozoic rocks are exposed in the Andean frontal monocline in a narrow belt that follows the regional northeast-southwest trend. Here the Cretaceous rocks were uplifted up to 10 km above the same units in the flanking foredeep basin. The exposed Cretaceous and Tertiary beds show moderate to steep northwestward dips. These dips are consistent over much of the study area. To the southeast the dip of the monocline steepens, and in some places the beds are overturned. The seismic data constrain the structural interpretation at depth.

Based on the seismic and surface geology information, the frontal monocline has been interpreted as the hanging wall forelimb of a major fault bend fold structure [Suppe, 1983]. A blind thrust fault is located in a bedding plane decollement beneath the mountain front monocline. Reflectors correlated to the La Luna Formation and older rocks are involved in the

hanging wall cutoffs. The thrust fault slips basinward into a bedding plane decollement in the Upper Cretaceous Colon-Mito Juan shales (Figures 8, 10, 12, 14, and 16). This interpretation requires the presence of a regional southeast dipping ramp south of the sections. The structural configuration of the frontal monocline is complicated by northwest dipping backthrusts and southwest dipping high-angle reverse faults.

In the section A-A' (Figures 6, 7, and 8), Upper Cretaceous and Tertiary units are folded in the La Friata anticline by a series of backthrusts. One backthrust fault is predicted in a bedding plane decollement in the Colon-Mito Juan formations stepping up into a higher detachment at the base of the Paleocene. Although prediction of this fault explains at least 2 km of shortening in the Lower Tertiary units, the major shortening in the anticline occurs within the Oligocene units. The Friata-1X well log shows an unusually thick sequence (1658 m) of Oligocene Leon Formation shales (Figures 7 and 8). The thick shale sequence is interpreted as a duplex consisting of five overlapping ramp anticlines. The solution shown in Figure 8 requires approximately 17 km of slip at the Oligocene/Eocene contact. If, however, the original thickness of the Oligocene Leon Formation in the line of the section was 500-800 m, only one or two overlapping ramp anticlines

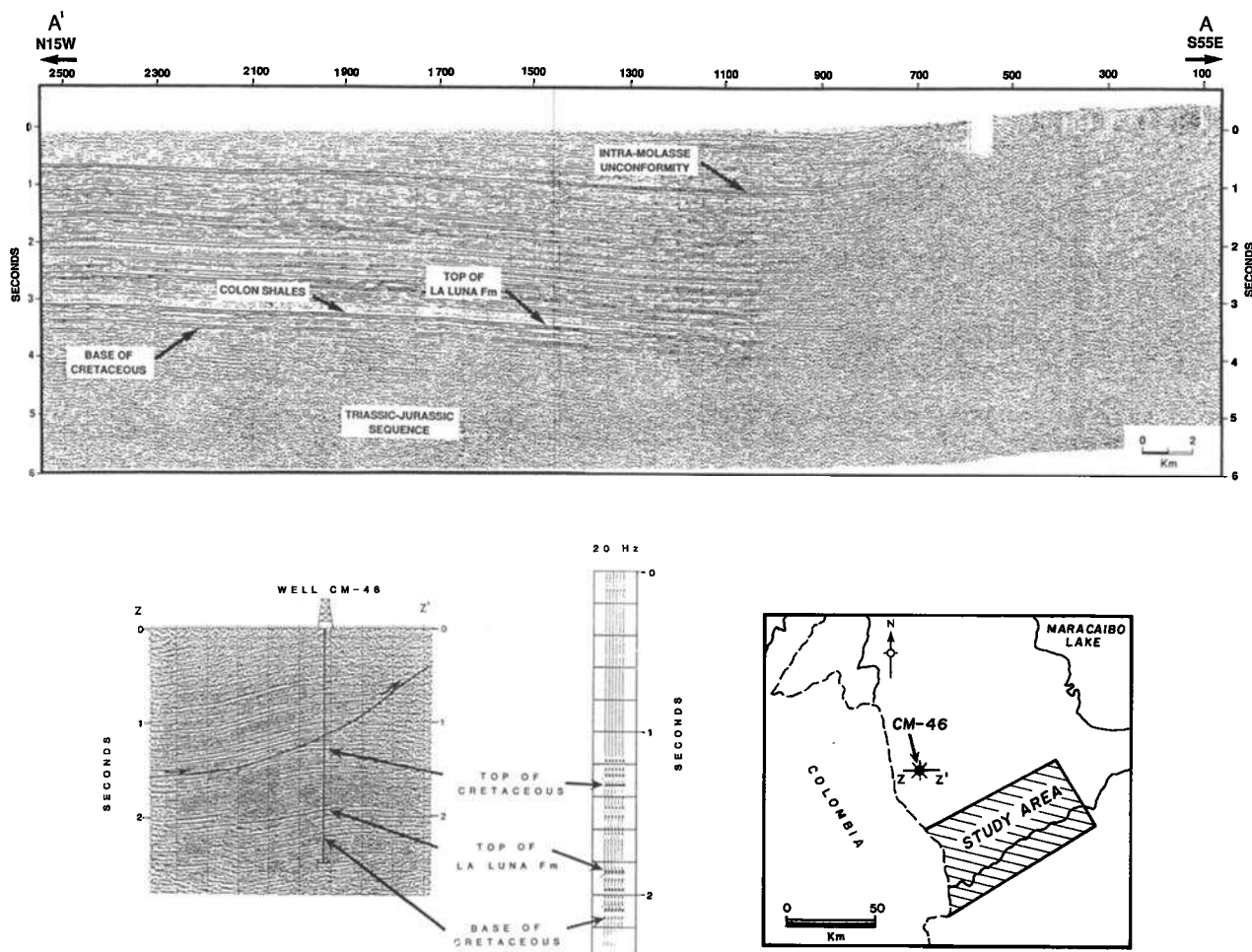


Fig. 5. Identification of the stratigraphic units in the foredeep basin (section A-A' in Figure 2) based on interpreted seismic line Z-Z' and synthetic trace of well CM-46 in the Los Manueles oil field.

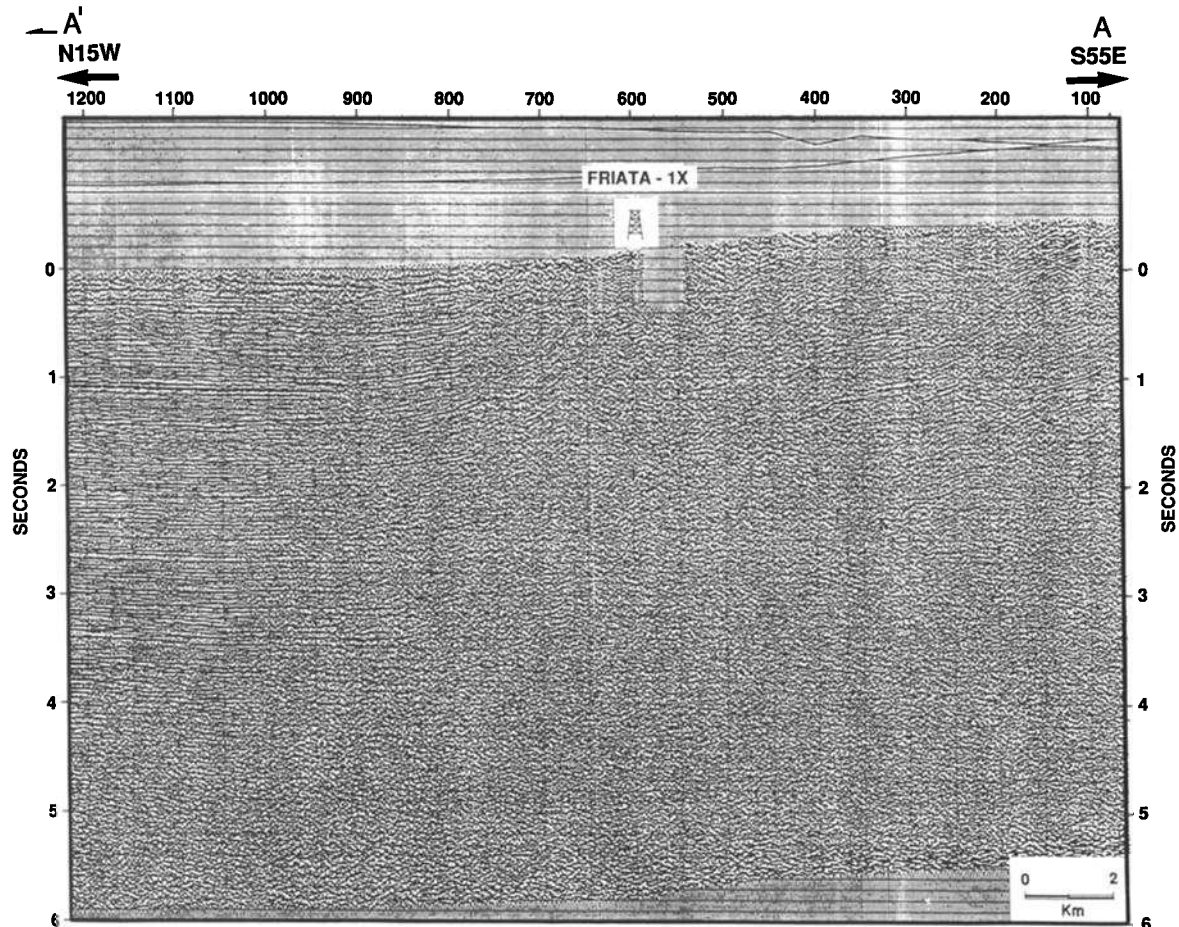


Fig. 6. Seismic section A-A' showing the location of the Friata-1X well.

and 4-8 km of slip would be required. Structural solutions with one or two ramps are compatible with the dip log data from the Friata-1X well (Figure 7), which suggests two or three distinct dip domains.

In seismic section B-B' (Figure 9), subhorizontal reflectors at 4.1-4.4 s underlie steeply dipping reflectors in the monocline. The subhorizontal reflectors are interpreted as undeformed Cretaceous units in the footwall (Figure 10) beneath a bedding plane decollement in the Upper Cretaceous Colon-Mito Juan shales.

Structural discontinuity of seismic reflections in the upper part of the profile C-C' (Figure 11) was related to a backthrust fault within the Miocene units (Figure 12). The Rio Escalante fault in section D-D' (Figures 13 and 14) and E-E' (Figures 15 and 16) is also a backthrust with 1-3 km of displacement. The Rio Escalante fault trace is oblique to the mountain front (Figure 2). In section D-D' this backthrust appears to step up from a decollement at the top of the Paleocene, whereas in section E-E' the fault is steeper and probably rooted in a bedding plane in Cretaceous units. Two more minor backthrusts are shown in Figure 16: one in the upper molasse sequence at the front of the monocline and another in the Cretaceous south of the Rio Escalante fault.

Two imbricated fault-related folds steepen the frontal monocline in section D-D' (Figures 13 and 14). Steepening of middle Miocene strata indicates post middle Miocene formation

of this duplex structure. Shallow onlapping reflections in the upper part of the profiles (Figures 13 and 15) suggest a period of Plio-Pleistocene erosion and monocline growth.

The fault bend fold interpretation for the Andean front would explain the 12 km of vertical uplift of the pre-Cretaceous units traversed by the sections. A minimum lateral shortening of 20 km is estimated for the interpreted structures.

#### *Foredeep Basin*

The foredeep basin is the structural depression of the southern Maracaibo Basin that is adjacent and parallel to the Venezuelan Andes. The basin is an asymmetric trough that contains up to 12 km of Cretaceous and Tertiary sediments. This sedimentary section is dominated by more than 6 km of Miocene to Recent molasse deposits derived from the erosion of the rising Andes. A major regional unconformity separates this thick clastic wedge into two different molasse units (Figures 5, 6, and 17). Structural deformation of the foredeep basin is mainly restricted to the mountain front; however, some minor normal faults offset seismic horizons correlated with the Cretaceous units (Figures 6, 8, and 15). The displacement of the faults is apparently absorbed or truncated in the reflection free zone produced by the Cretaceous Colon-Mito Juan formation shales.

We attribute the structural configuration of the foredeep and

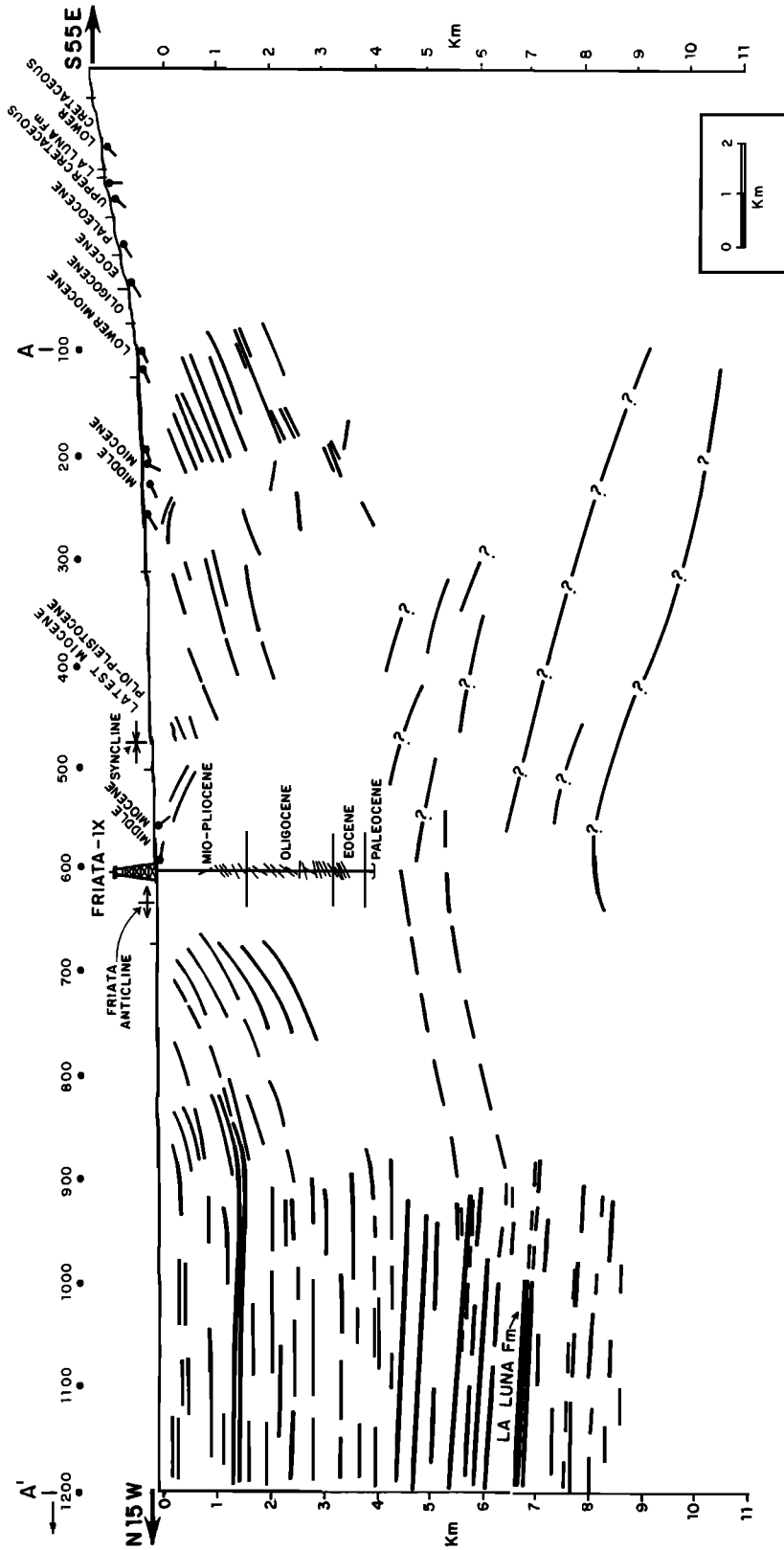


Fig. 7. Section A-A' converted to true depth with no vertical exaggeration, showing seismic reflectors and apparent bedding dips from the Friata-IX well and surface measurements.

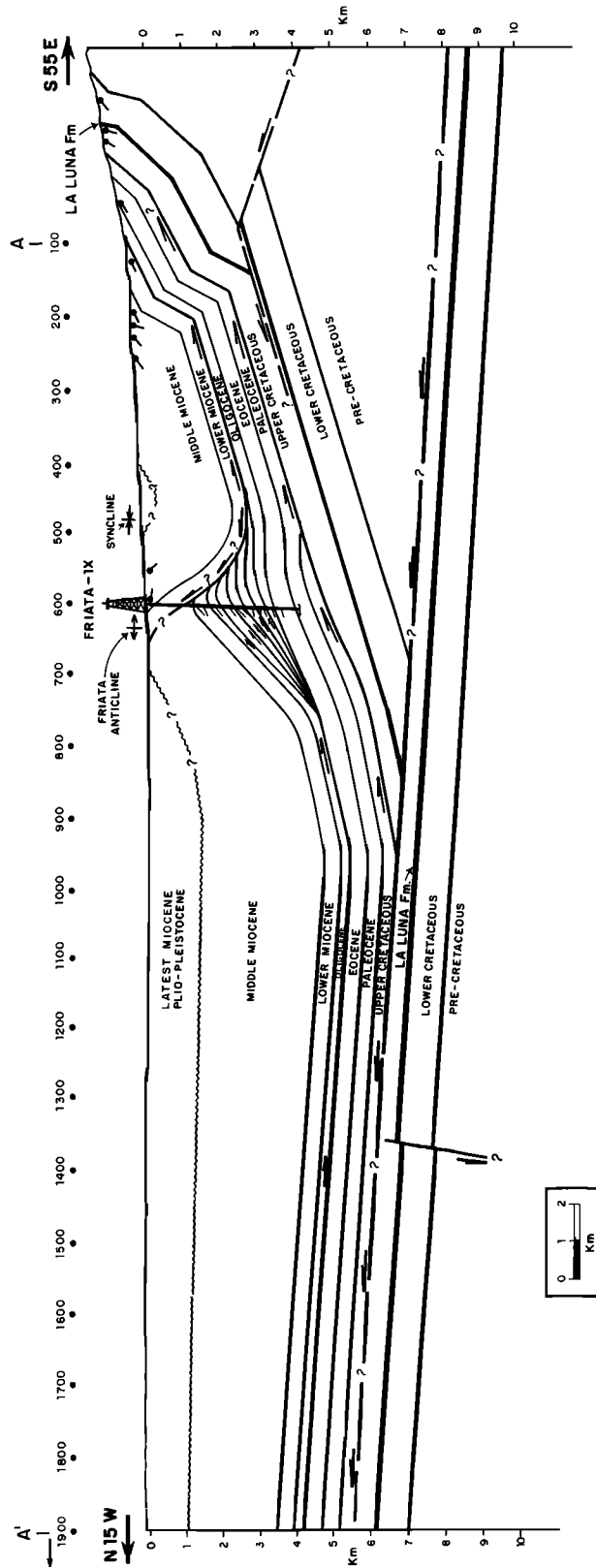


Fig. 8. Structural interpretation of section A-A'. Upper Cretaceous and Tertiary units are folded in the La Friata anticline by a series of backthrusts. One backthrust fault is predicted in a bedding plane decollement in Upper Cretaceous shales. The thick Oligocene shale sequence penetrated by Friata-IX is interpreted as a duplex consisting of five overlapping ramp anticlines. See text for discussion of an alternative solution consisting of only one or two overlapping ramp anticlines.



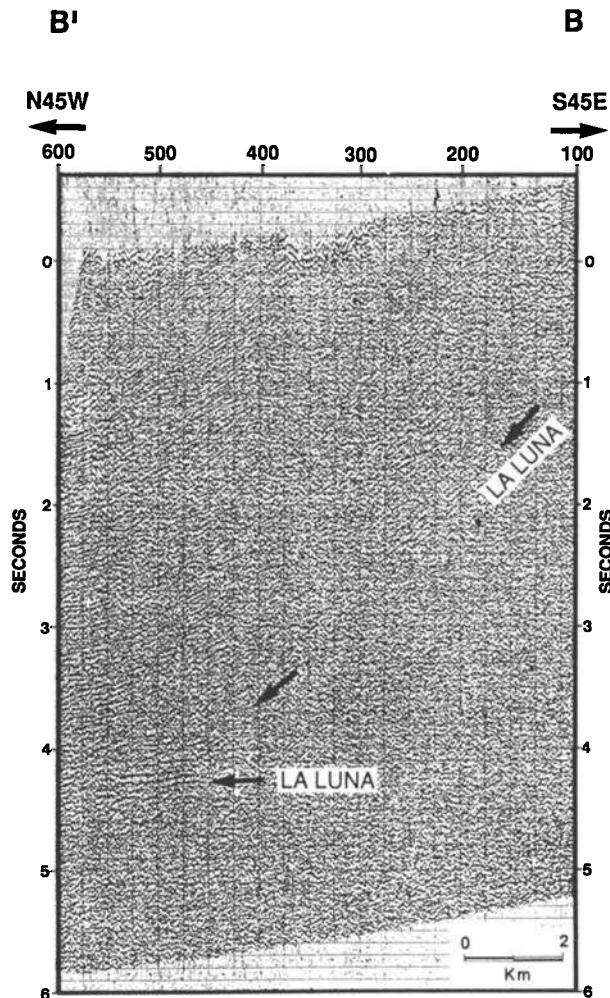


Fig. 9. Seismic section B-B'. Beneath the steeply dipping reflectors in the monocline note the subhorizontal reflectors at 4.1-4.4 s. Dashed lines indicate the location of reflectors interpreted as being associated with the La Luna Formation.

the regional southeastward tilting of the seismic reflectors to flexure of the lithosphere as a response to tectonic loading by the Venezuelan Andes. The normal faults at the base of the sedimentary section in the foredeep basin may be secondary features produced by the bending stresses.

## DISCUSSION

### *Crustal Scale Fault-Related Folding and Wedge Structure*

The seismic sections presented in this paper show considerable shortening by imbricate backthrust and breakthrough faults within the monocline on the northwestern front of the Venezuelan Andes. Sediments monoclinaly draped over a vertical or normal basement fault would be extended by normal faulting or ductile thinning [Brown, 1988] which is not observed. The reverse faulting observed in the sediments therefore indicates reverse faulting and shortening in the basement. We have therefore interpreted the Venezuelan Andes as a northwestward verging fault bend fold structure

deforming over a crustal-scale ramp (Figures 18a and 18b). Estimates of the fault dip range from 20°-25° based on gravity models [Kellogg and Bonini, 1982, 1985] to about 27° from fault bend fold theory for a 38° frontal limb dip [Suppe, 1983]. The simplified solution shown in Figure 18b shows a low-angle (20°) ramp, (1) from fault bend fold theory for the 20°-23° frontal limb dips in profiles A, C, and D (Figures 8, 12, and 14) and (2) so that the monocline on the southeastern flank of the Venezuelan Andes could be explained as the backlimb of the same fault bend fold that formed the northwestern flank. The low-angle ramp lies within the range of solutions compatible with the gravity data [Kellogg and Bonini, 1982, 1985]. The high dip angles shown in profiles B, C, D, and E (Figures 10, 12, 14, and 16) are interpreted as frontal limb dips produced by imbricate thrust faults not shown in Figure 18b. The Andes are thrust 20-30 km northwestward toward the Maracaibo Basin. Vertical relief on the top of basement is more than 12 km. The fault ramps up to a decollement horizon within the Cretaceous Colon-Mito Juan formation shales. The northwestward movement of the Andean basement folded the hanging wall sedimentary sequence and produced the steep monocline of the north Andean flank. Deep subhorizontal reflectors beneath the monocline are interpreted as undeformed Cretaceous units in the footwall (Figures 9 and 10). Backthrusts in the monocline produce a wedge geometry [Tearpock and Bischke, 1991] and reduce the amount of blind slip required on the decollement northwest of the Venezuelan Andes.

Our interpretation is supported by new seismic lines [Audemard, 1991] that show a major decollement in the Upper Cretaceous Colon shales and backthrusting within Colon, Lower Tertiary, and Molasse units. Backthrusts consist of fault bend folds, fault propagation folds, and duplexes. Northwest verging duplexes involve the basement.

Nonscaled rock models of low-angle reverse faults deformed under confining pressure show late stage basement structures developed through fracture, imbricate faulting, and rotation of basement blocks between zones of cataclasis [Chester et al., 1988]. Overlying the uplifted basement block in the models is a continuous, flat-topped, asymmetric kink fold in the veneer having one near-vertical limb. The steeply dipping limb is analogous to the monocline on the northwestern flank of the Venezuelan Andes.

### *Timing of Deformation*

The thick molasse deposits accumulated in the foredeep basin provide the most direct evidence for timing the Andean thrusting and uplift. These sediments, represented by the El Guayabo Group (El Palmar, Isnotu, and Betijoque formations), suggest that the Venezuelan Andes started to rise as early as the early Miocene, with the formation of the monocline and the greatest uplift in late Miocene to Pleistocene times.

The El Guayabo Group has always been described as one uninterrupted cycle of continental sedimentation [Gonzalez de Juana et al., 1980]. However, the seismic data show an important intramolasse unconformity, indicating that two different sedimentation cycles are separated by a period of tectonic interruption. This unconformity is most pronounced in the western portion of the study area (Figure 17). In outcrops of the north Andean mountain front, the intramolasse discontinuity is not readily observed; however, in the southwestern flank (Tachira Depression) the unconformity can

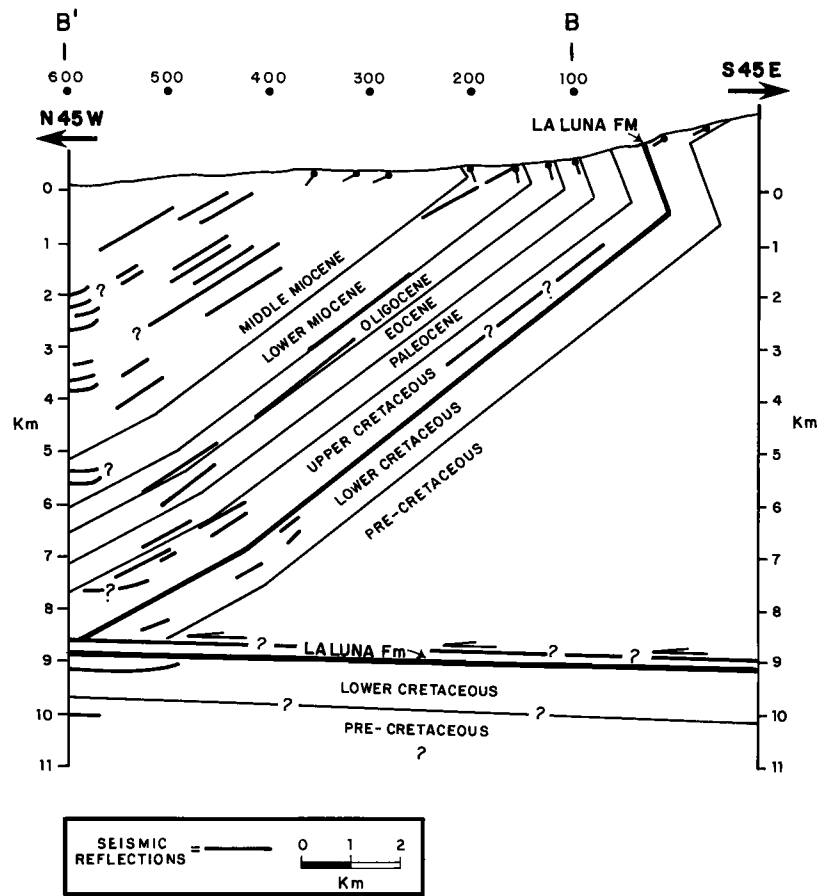


Fig. 10. Structural interpretation of seismic section B-B' with no vertical exaggeration, showing seismic reflectors and apparent bedding dips at the surface. The subhorizontal reflectors are interpreted as undeformed Cretaceous units in the footwall beneath a bedding plane decollement in the Upper Cretaceous shales.

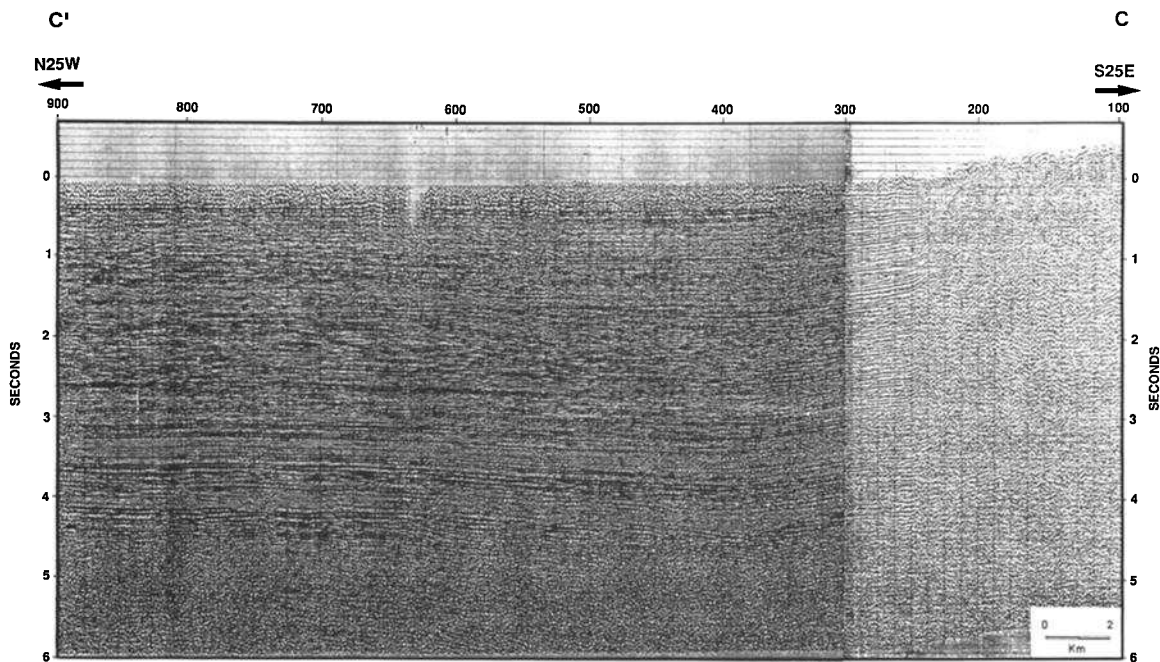


Fig. 11. Seismic section C-C'.

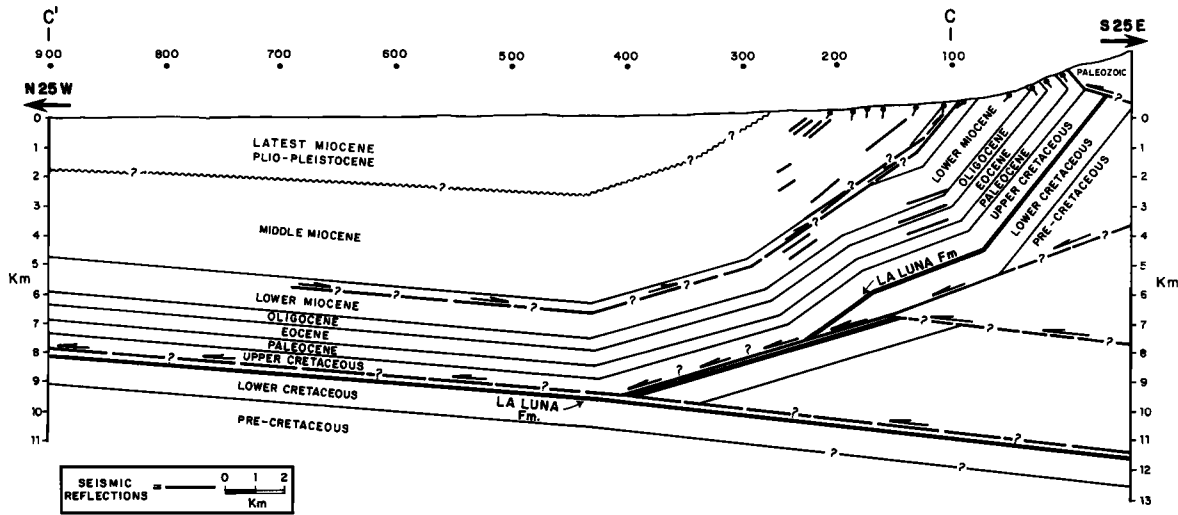


Fig. 12. Structural interpretation of section C-C' showing seismic reflectors and apparent bedding dips from surface measurements.

be distinguished and occurs beneath the fluvial latest Miocene sediments of the La Cope Formation [Macellari, 1984]. This intra-Miocene event has also been found in seismic sections of the Barinas-Apure Basin and in most of the intermontane basins of northwestern Colombia [Meier et al., 1987; Chigne, 1985; Eva et al., 1989].

**SUMMARY**

The seismic reflection profiles of the northwestern flank of the Venezuelan Andes presented in this paper are consistent

with our interpretation of the range as a fault-related fold structure deformed over a crustal-scale ramp. The ramp dips about 20°-30° southeast, and the Andean basement was thrust 20-30 km northwestward toward the Maracaibo Basin. The blind thrust fault steps into a bedding plane decollement horizon within the Cretaceous Colon-Mito Juan formation shales. The northwestward movement of the Andean basement folded the hanging wall sedimentary sequence and produced a steep monocline. Backthrusts in the monocline produce a

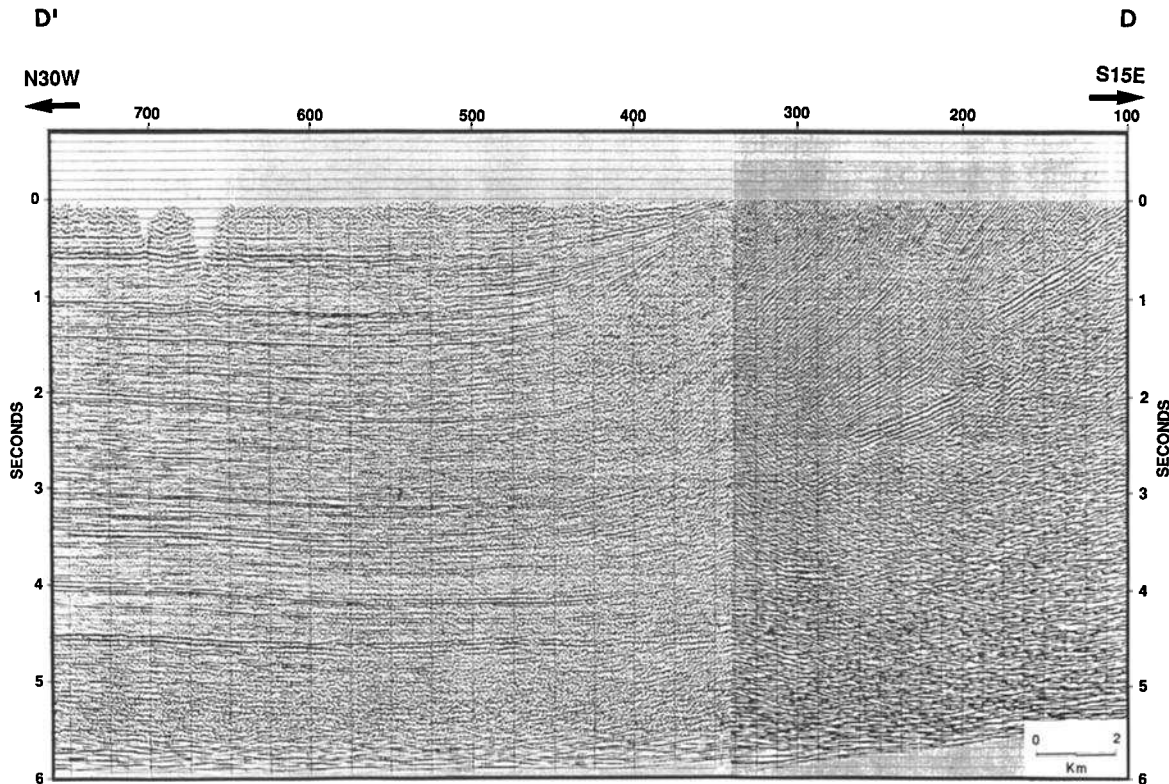


Fig. 13. Seismic section D-D'.

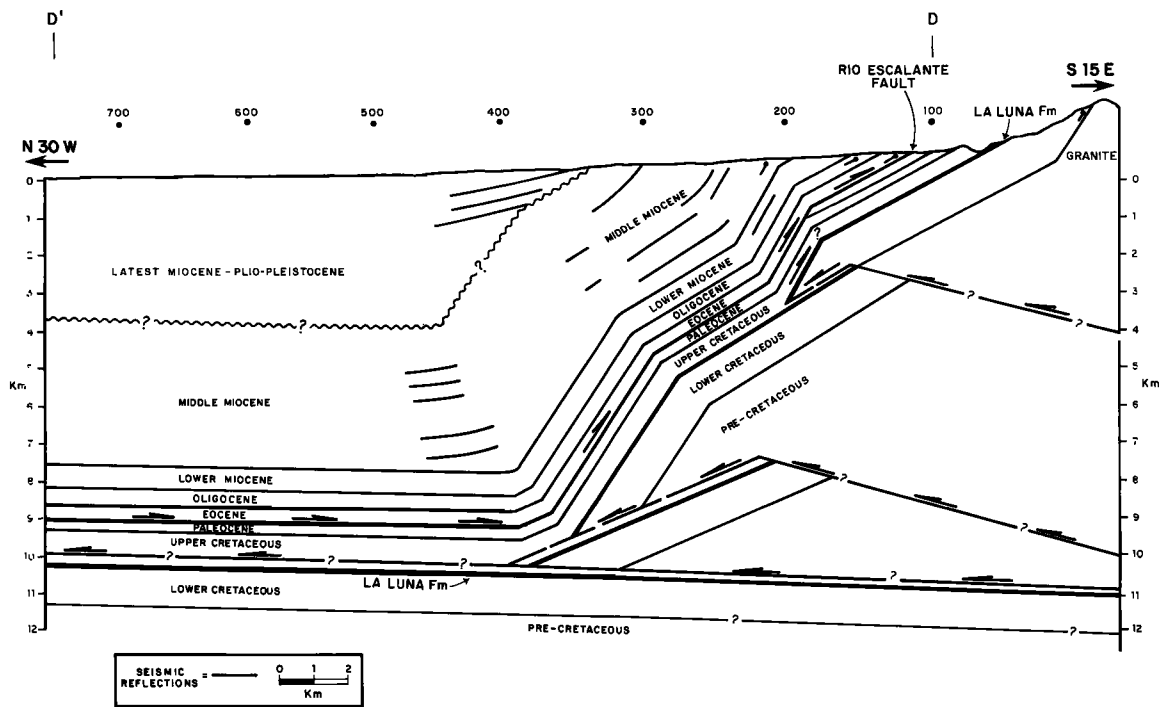


Fig. 14. Structural interpretation of seismic section D-D' showing seismic reflectors and apparent bedding dips. Imbricate southeast dipping thrusts within the basement and northwest dipping backthrusts within the sedimentary cover steepen the monocline. Discordant reflectors between SP350 and SP450 suggest a late Miocene unconformity associated with the formation of the monocline.

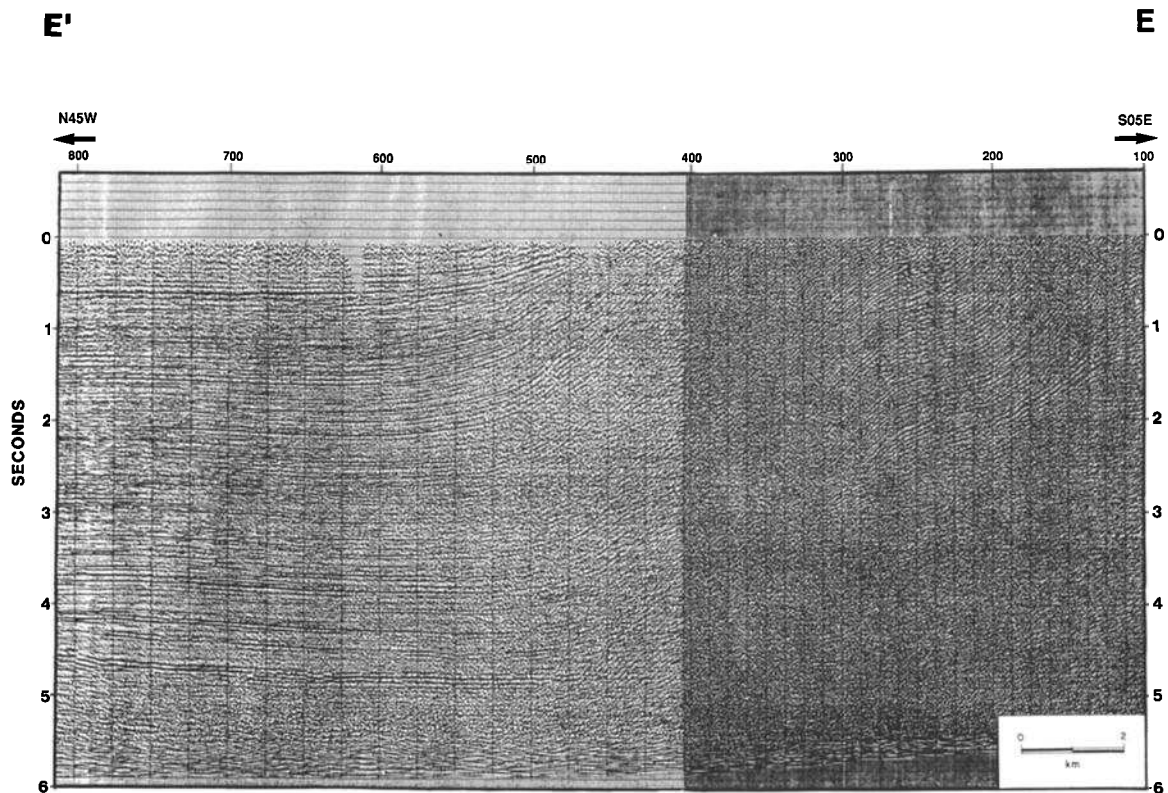


Fig. 15. Seismic section E-E'.

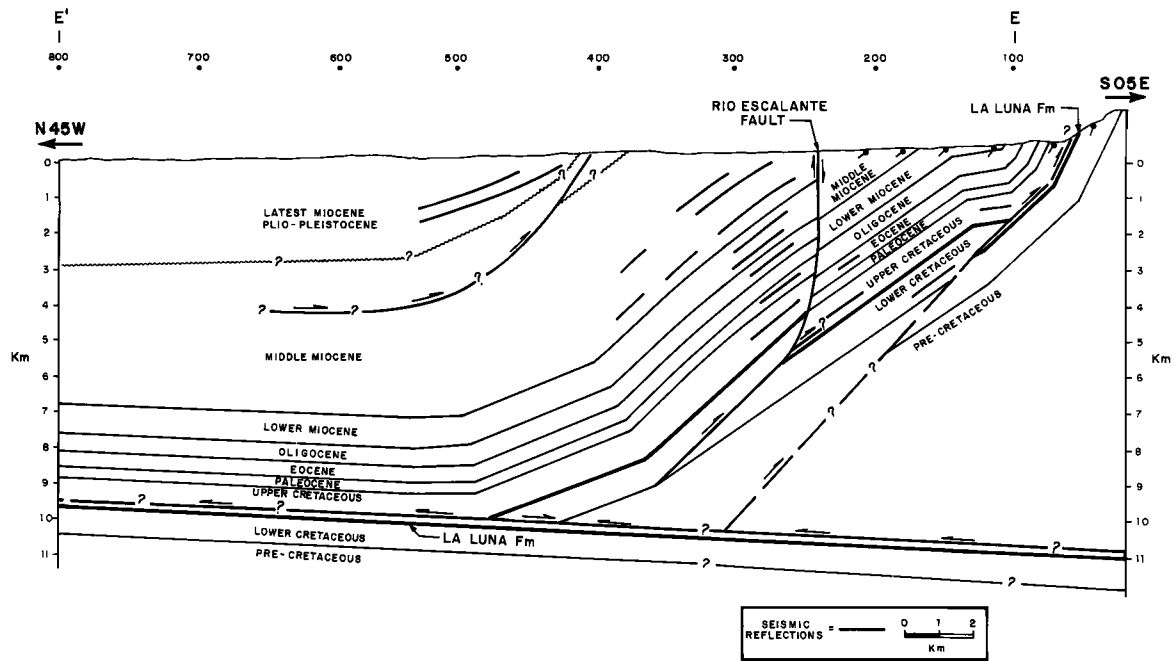


Fig. 16. Structural interpretation of section E-E'. The Rio Escalante fault is a backthrust with a trace oblique to the mountain front (Figure 2). The discordance of shallow reflectors at SP500 is interpreted as being produced by growth of the monocline.

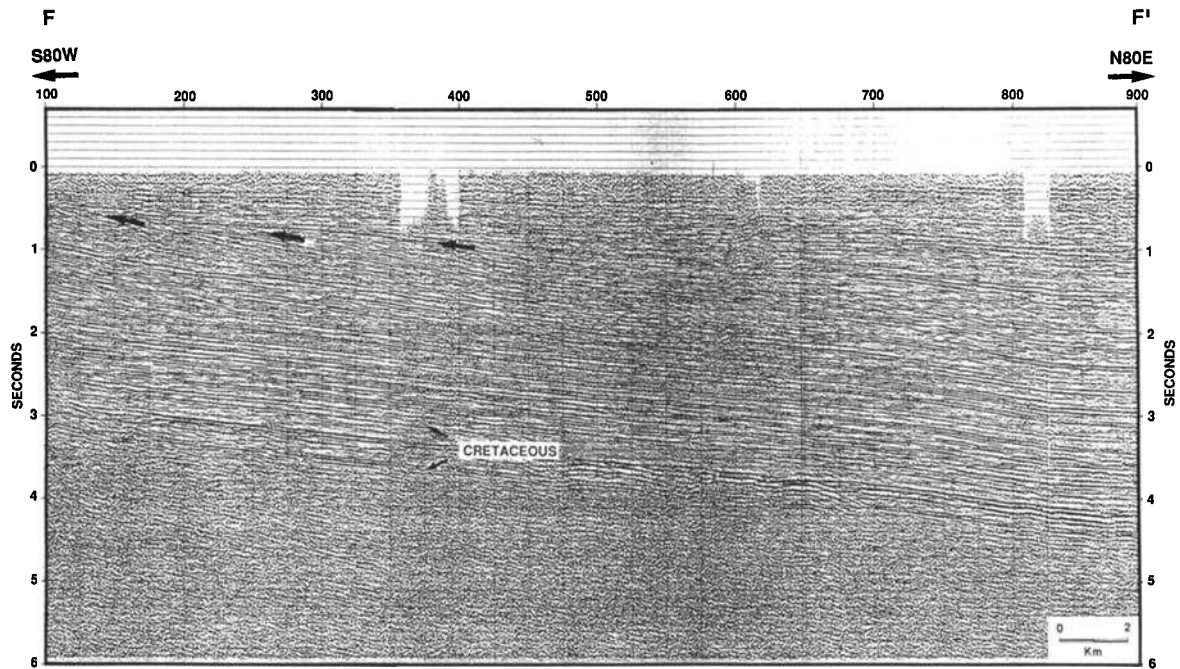


Fig. 17. Seismic line F-F' showing truncation of reflections (solid arrows) beneath the intramontane unconformity. This unconformity is most pronounced in the western portion of the study area.

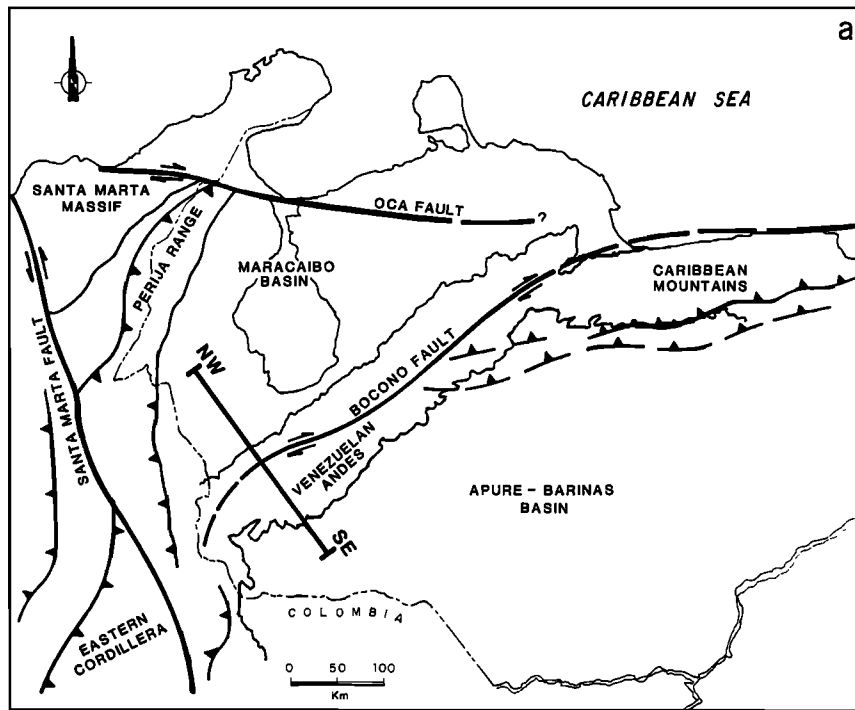


Fig. 18a. Tectonic location map for schematic cross section of the Venezuelan Andes (Figure 18b).

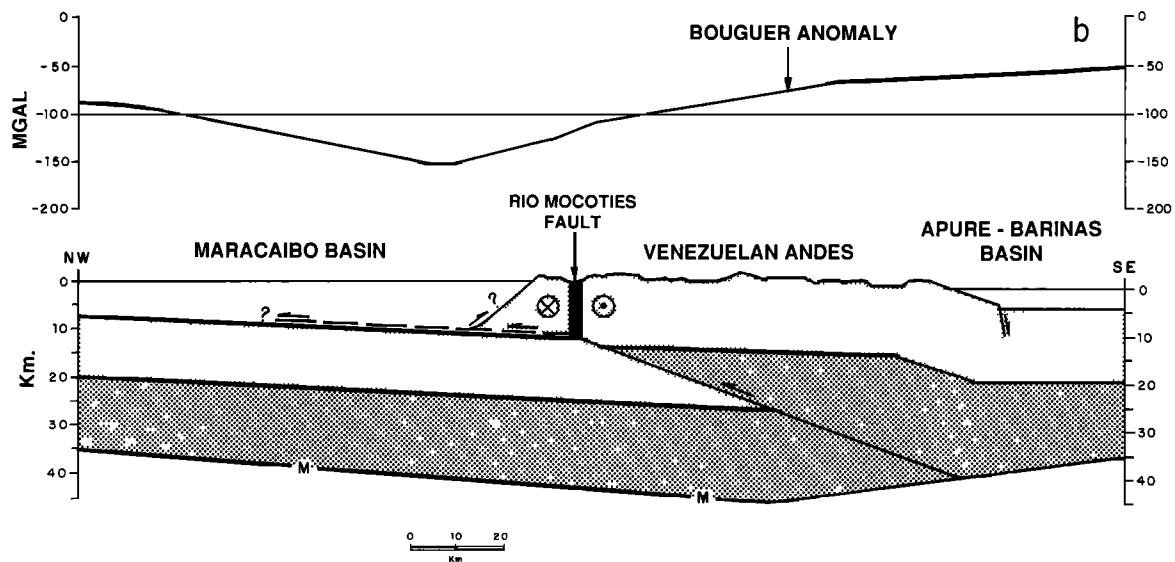


Fig. 18b. Schematic cross section of the Venezuelan Andes located in Figure 18a. We have interpreted the Venezuelan Andes as a fault bend fold structure deforming over a crustal-scale southeast dipping ramp. Estimates of the fault dip range from  $20^{\circ}$ - $25^{\circ}$  based on gravity models [Kellogg and Bonini, 1982, 1985] to about  $27^{\circ}$  from fault bend theory for a  $38^{\circ}$  frontal limb dip [Suppe, 1983]. In the solution shown ( $20^{\circ}$  ramp) the monocline on the southeastern flank of the Venezuelan Andes is the backlimb of the fault bend fold. The Andes are thrust 20-30 km northwestward toward the Maracaibo Basin. Vertical relief on the top of basement is more than 12 km. Backthrusts in the monocline produce a wedge geometry and reduce the amount of blind slip required on the decollement northwest of the Venezuelan Andes.

wedge geometry and reduce the amount of blind slip required on the decollement northwest of the Venezuelan Andes. Rock models of low-angle reverse faults deformed under confining pressure show late stage basement structures similar in shape to the predicted fault bend fold.

The monocline on the northwestern flank of the Venezuelan Andes is the result of crustal shortening in a northwest-southeast direction. This shortening has been attributed to the convergence of the Caribbean plate and the North Andean microplate at the Caribbean marginal fault.

Two molasse cycles separated by an intra-Miocene discontinuity are present in the southwestern part of the foredeep basin. The first cycle corresponds to the Miocene El Palmar and Isnotu formations. The second cycle corresponds to the late Miocene-Plio-Pleistocene Betijoque Formation. These thick (up to 8 km) continental molasse deposits indicate that the Andean uplift had begun as early as early Miocene,

although the formation of the monocline and the greatest uplift occurred in late Miocene to Pleistocene time. The intramolasse unconformity may be related to the initial late Miocene collision of the Panama volcanic arc with northwestern South America.

*Acknowledgments.* Christopher Kendall advised and directed all the stratigraphic aspects of the study. We thank Reginald Shagam and Wayne Narr for critical reviews. Maraven, S.A., Venezuela, generously provided all the seismic information. Albert Bally kindly provided a copy of Felipe Audemard's Ph.D. thesis. The second author was supported in part by a grant from the Petroleum Research Fund of the American Chemical Society. We want to thank INTEVEP, S.A., Venezuela, for supporting research by the first author at the University of South Carolina and for authorizing the publication of this paper.

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- (Received May 28, 1992;  
revised June 29, 1993;  
accepted July 13, 1993.)